Lecture 3: Boulder Summer School Steve Tobias (University of Leeds)

A mean field model of the Solar Dynamo



Dynamics vs Statistics

"The climate is what you expect; the weather is what you get." *Attributed to Mark Twain*.

"More than any other theoretical procedure, numerical integration is also subject to the criticism that it yields little insight into the problem. The computed numbers are not only processed like data but they look like data, and a study of them may be no more enlightening than a study of real meteorological observations. An alternative procedure which does not suffer this disadvantage consists of deriving a new system of equations whose unknowns are the statistics themselves....

...This procedure can be very effective for problems where the original equations are linear, but, in the case of non-linear equations, the new system will inevitably contain more unknowns than equations, and can therefore not be solved, unless additional postulates are introduced."

Edward Lorenz, The Nature and Theory of the General Circulation of the Atmosphere (1967)

"Direct Statistical Simulation" (DSS)

Direct Statistical Simulation: Why *simulate* the statistics?

Low-order statistics: smoother in space than instantaneous flow.

Statistics evolve slowly in time, or not at all, and hence may be described by a fixed point, or at least a slow manifold. Statistics usually describe mean/average behaviour and variations about that mean (e.g. 2pt correlations) Statistics are less sensitive to changes in underlying parameters than detailed dynamics.

In geophysics/astrophysics correlations are *non-local* and highly anisotropic and inhomogeneous. Statistical formulations *must* respect this. They should also respect conservation laws

Solution of Statistical Equations is an old idea: Bousssinesq, Reynolds, Lorenz, Herring, Kraichnan, Frisch, Salmon, Monin & Yaglom, Majda

Direct Statistical Simulation: the options...

- For any dynamical system one could try to simulate pdfs
 - Perron-Frobenius Theory (see e.g. Beck & Schloegl 1993)
 - Fokker-Planck/Liouville equations (see e.g. Kadanoff 2000, Alwalla & Marston 2016, Cho et al 2016, Chen & Majda 2017)
- Kolmogorov/Kraichnan moment hierarchies
 - Usually assume isotropy/homogeneity of statistics
- Large Deviation Theory
 - Extreme/rare events or flipping (see e.g. Bouchet & Simonet 2009, Laurie & Bouchet 2015)
- Use cumulants
 - take into account inhomogeneity/anisotropy
 - Statistical closures
 - Relationship with deterministic approximations?



See e.g. Foias, Manley & Temam 2001

$$\dot{x} = \sigma(y - x) + \eta_1(t)$$

$$\dot{y} = x(\rho - z) - y + \eta_2(t)$$

$$\dot{z} = xy - \beta z + \eta_3(t)$$

• What does pdf look like?



https://hypertextbook.com/chaos/strange/

$$\dot{x} = \sigma(y - x) + \eta_1(t)$$
$$\dot{y} = x(\rho - z) - y + \eta_2(t)$$
$$\dot{z} = xy - \beta z + \eta_3(t)$$

• Simulation of pdfs: L/FPE

$$\frac{\partial P(\vec{x},t)}{\partial t} = -\hat{L}_{FPE}P(\vec{x},t)$$



 $\hat{L}_{FPE}P = \vec{\nabla} \cdot \left[(\sigma(y-x), \ x(\rho-z) - y, \ xy - \beta z)P \right] - \Gamma \nabla^2 P$

Deterministic flow

noise

Use cumulant expansions

$$\dot{q}_{i} = F_{i} + L_{ij} \ q_{j} + Q_{ijk} \ q_{j}q_{k} + \eta_{i}.$$

$$\begin{array}{c} \text{ODE}\\ \text{SYSTEM} \end{array}$$

$$q_{i} = \overline{q_{i}} + q_{i}' \qquad \begin{array}{c} \text{REYNOLDS}\\ \text{DECOMPOSITION} \end{array}$$

$$c_{i} \equiv \overline{q_{i}}\\ c_{ij} \equiv \overline{q_{i}'q_{j}'} \qquad \begin{array}{c} \text{DEFINE}\\ \text{CUMULANTS} \end{array}$$

$$c_{ijk} \equiv \overline{q_{i}'q_{j}'q_{k}'}\\ c_{ijk\ell} \equiv \overline{q_{i}'q_{j}'q_{k}'q_{\ell}'} - c_{ij} \ c_{k\ell} - c_{ik} \ c_{j\ell} - c_{i\ell} \ c_{jk} \end{array}$$

• Derive cumulant system using Hopf Functionals (or brute force)

$$\dot{q}_i = F_i + L_{ij} \ q_j + Q_{ijk} \ q_j q_k + \eta_i.$$

$$q_i = \overline{q_i} + q'_i$$

1st Cumulant Equation

$$\frac{dc_i}{dt} = \frac{\overline{dq_i}}{dt}$$
$$= F_i + L_{ij} c_j + Q_{ijk} (c_j c_k + c_{jk})$$

$$\frac{dc_{ij}}{dt} = 2\left\{ \overline{\frac{dq'_i}{dt}} q'_j \right\}$$
$$= \left\{ 2L_{ik} c_{kj} + Q_{ik\ell} \left(4c_k c_{\ell j} + 2c_{k\ell j} \right) \right\} + 2\Gamma_{ij}$$

Li et al (2022)

2nd Cumulant Equation

$$\begin{split} \dot{q}_{i} &= F_{i} + L_{ij} \ q_{j} + Q_{ijk} \ q_{j}q_{k} + \eta_{i}. \\ \hline \frac{dc_{ijk}}{dt} &= 3 \left\{ \frac{dq'_{i}}{dt} q'_{j}q'_{k} \right\} \\ &= \{3L_{im} \ c_{mjk} + 6Q_{imn} \ (c_{m}c_{njk} + c_{mj}c_{nk})\} - \frac{c_{ijk}}{\tau_{d}} \\ \hline \\ & (a) \\ \hline \\ (b) \\ (c) \\ (c)$$

0

0

z 20

0

J 20 10 0 x -10

 $\overline{x'z'}$

 $O(10^{-5})$

Cumulant Expansion for PDEs





DSS: Cumulant Expansion to 2nd Order (CE2)

"The discarding of the fluctuating self-interaction then corresponds to closing the system of moment equations by discarding the third order cumulants" (Herring 1963).

$$egin{aligned} \mathbf{c}(\mathbf{r}) &\equiv \overline{\mathbf{u}(\mathbf{r})}, \ \mathbf{c}(\mathbf{r}_1,\mathbf{r}_2) &\equiv \overline{\mathbf{u}'(\mathbf{r}_1)\otimes\mathbf{u}'(\mathbf{r}_2)} \end{aligned}$$





Wave — Mean-Flow Interaction

Reynolds Forcing of Mean-Flow

$$\partial_t \mathbf{c}(\mathbf{r}_1) = \mathcal{L}[\mathbf{c}(\mathbf{r}_1)] + \mathcal{N}[\mathbf{c}(\mathbf{r}_1), \ \mathbf{c}(\mathbf{r}_1)] + \int \mathcal{N}[\mathbf{c}(\mathbf{r}_1, \mathbf{r}_2), \delta(\mathbf{r}_1 - \mathbf{r}_2)\mathbf{I}] d\mathbf{r}_2$$

$$\partial_t \mathbf{c}(\mathbf{r}_1,\mathbf{r}_2) \,=\, \mathcal{L}_{\mathbf{c}(\mathbf{r}_1)}[\mathbf{c}(\mathbf{r}_1,\mathbf{r}_2)] + \mathcal{L}_{\mathbf{c}(\mathbf{r}_2)}[\mathbf{c}(\mathbf{r}_1,\mathbf{r}_2)]$$

Paradigm Problem:

Barotropic turbulence. Two-dimensional hydrodynamics on a sphere

$$\partial_t q = \mathcal{L}[q] + \mathcal{N}[q \ q] + f(t)$$

$$\partial_t \zeta = -\kappa \zeta - \nu_4 \nabla^4 \zeta - \mathbf{v} \cdot \nabla (\zeta + 2\Omega \sin \theta) + \eta(t)$$

$$\zeta = (\nabla \times \mathbf{v}) \cdot \hat{\mathbf{r}}$$

$$\zeta(\theta, \phi) = \sum_{\ell=0}^L \sum_{m=-\ell}^\ell \zeta_{\ell,m} Y_\ell^m(\theta, \phi) \qquad \begin{array}{l} \text{Radial component} \\ \text{of the vorticity} \end{array}$$

Other problems at CE2:

2D turbulence on beta-plane (Farrell & Ioannou 2007, Tobias & Marston 2013, Bakas & Ioannou 2014) Joint instabilities in the Tachocline (Plummer et al 2017) MRI in shearing box (Squire & Bhattacherjee 2014) Convection in a Busse Annulus Saturation of Inertial (symmetric) instabilities

Comparison

- Compare two techniques:
- DNS using spherical geodesic grid
 - Second-order leapfrog
 - Robert filter
 - Multigrid algorithm
- Fully spectral Direct Statistical Simulation
 - Semi-implicit Krylov method
 - Approach to a simple attractor (fixed point, limit cycle?)

What do we expect?

- Small-scale turbulence will drive jets via PV homogenisation (Rhines 1975, McIntrye 2003, Dritschel & McIntyre 2008, Vallis & Maltrud 1993, Manfroi & Young 1998, Sukoriansky & Galperin, 2008
 - Correlations of nonlinear Rossby waves drives zonal flow via Reynolds Stresses
 - Non-trivial interactions lead to generations of mean flows and angular momentum transport

DNS Movie of hydro jet formation



Relative Vorticity

Hydrodynamics: DNS



Zonally averaged zonal radial vorticity



Zonally averaged zonal azimuthal velocity

DSS Movie of hydro jet formation



Relative Vorticity

Hydrodynamic: DSS

2000



1000

DAYS

2000.5 days

60N-

30N-

EQ

30S ·

60S ·

DSS reproduces driving of jets.

This is possible even for a cumulant expansion truncated at second order.

These jets are therefore driven directly via Reynolds stresses.

Recall: No inverse (or forward) cascade in this model.

Hydrodynamic: DSS



- Detailed comparison: DSS reproduces shape of jet very well (except at poles where mean is small)
- No eddy-eddy scattering

Covariance/2pt correlation?



Rank Instability of CE2

Nivarti, Kerswell, Marston & Tobias arxiv:2202.04127

 $\dot{x} = 2x$ $\dot{y} = -y$

 $rac{d}{dt} \ \overline{xx} = 4 \ \overline{xx}$ $rac{d}{dt} \ \overline{xy} = \overline{xy}$ $rac{d}{dt} \ \overline{xx} = -2 \ \overline{yy}$



How to improve on QL/CE₂/SSST?

- Generalise the quasilinear approximation
 - Derive statistical theory (GCE2) corresponding to GQL approximation [Nivarti, Marston & Tobias in preparation]
- Include eddy/eddy → eddy interactions
 CE2.5/CE3 [Marston et al (2015) arXiv:1412.0381]
- Change the definition of averaging to ensemble averaging [Allawala et al 2020]

How to improve on QL/CE₂/SSST?

$$rac{1}{ au} \ \mathbf{c}(\mathbf{r}_1,\mathbf{r}_2,\mathbf{r}_3) = \{\mathcal{N}[\mathbf{c}(\mathbf{r}_1,\mathbf{r}_2),\mathbf{c}(\mathbf{r}_1,\mathbf{r}_3)]\}$$

- CE2.5 is the generalization of the eddy-damped quasi-normal Markovian approximation (EDQNM) to anisotropic & inhomogeneous turbulent flows. Appears to be realizable.
- [Marston et al (2015) arXiv:1412.0381]

Improvements on CE2/ Beyond Quasilinearity Include eddy/eddy → eddy interactions (CE2.5/CE3)

 $\dot{c}_{1} = \mathcal{L}[c_{1}] + \mathcal{N}[c_{1}c_{1} + c_{2}]$ $\dot{c}_{2} = \mathcal{L}[c_{2}] + \mathcal{N}[c_{1}c_{2} + c_{3}] + \Gamma$ $\dot{c}_{3} = \mathcal{L}[c_{3}] + \mathcal{N}[c_{1}c_{3} + c_{2}c_{2} + c_{4}]$



Generalise the definition of mean to large scales (low modes) (GCE₂)

 $\begin{aligned} \partial_t q &= \mathcal{L}[q] + \mathcal{N}[q \ q] \quad q = \ell + h \\ \partial_t l &= \mathcal{L}[l] + \mathcal{N}[l \ l] + \mathcal{N}[h \ h] \\ \partial_t h &= \mathcal{L}[h] + \mathcal{N}[l \ h] \\ \partial_t (h, h) &= \mathcal{L}[(h, h)] + 2\mathcal{N}[l \ (h], h) \end{aligned}$

Statistical implementation of GQL



Summary



500.0 days

Barotropic











Model Reduction via POD

• Can we do the calculation keeping only the information that is needed? i.e. reduce the basis... $c_2 = \langle q'(\mathbf{x_1})q'(\mathbf{x_2}) \rangle$

$$c_2(\mathbf{x}_1, \mathbf{x}_2) = \sum_N \lambda_i \phi_i(\mathbf{x}_1) \phi_i(\mathbf{x}_2)$$

 $\approx \sum_{i=1}^N \lambda_i \phi_i(\mathbf{x}_1) \phi_i(\mathbf{x}_2)$
Schmidt Decomposition

Proper Orthogonal Decomposition Principal Component Analysis (see e.g. Holmes et al *Turbulence, Coherent Structures, Dynamical Systems and Symmetry, 2012*)





DSS as a subgrid model (Kuan Li)



very fine grid

DNS on a coarse grid + DSS of unresolved scales

- Can DSS be used dynamically?
- Think of the first cumulant as the DNS solution on a coarse grid
- Second cumulant is correlation of fluctuations from all the unresolved scales, projected onto the coarse grid
- Very expensive??
 - Just keep first PoD mode of second cumulant (how accurate?)
 - Matrix-free method (Li)

DSS as a subgrid model (Kuan Li)

$$\frac{\partial}{\partial t} \mathbf{f} = \mathcal{N}(\mathbf{f}, \mathbf{f}) + \mathcal{L}(\mathbf{f})$$
$$\mathbf{f} = \zeta \oplus n, \qquad \mathcal{N}(\mathbf{f}, \mathbf{f}) = \mathcal{J}(\psi, \zeta) \oplus \mathcal{J}(\psi, n)$$
$$C_{\mathbf{ff}} = \sum_{i} \mathbf{s}_{i} \otimes \mathbf{s}_{i},$$
$$\frac{\partial}{\partial t} C_{\mathbf{f}} = \mathcal{N}(C_{\mathbf{f}}, C_{\mathbf{f}}) + \mathcal{L}(C_{\mathbf{f}}) + \mathbf{h},$$
$$\frac{\partial}{\partial t} C_{\mathbf{ff}} = \sum_{i} \mathbf{s}_{i} \otimes \mathbf{p}_{i} + \mathbf{p}_{i} \otimes \mathbf{s}_{i}.$$
$$\mathbf{h} = \sum_{i} \mathcal{N}(\mathbf{s}_{i}, \mathbf{s}_{i})$$

 $\mathbf{p}_{i} = \left[\mathcal{N}(C_{\mathbf{f}}, \mathbf{s}_{i}) + \mathcal{N}(\mathbf{s}_{i}, C_{\mathbf{f}}) + \mathcal{L}(\mathbf{s}_{i})\right] + 2\tau_{d} \left[\mathcal{N}(\mathbf{h}, \mathbf{s}_{i}) + \mathcal{N}(\mathbf{s}_{i}, \mathbf{h})\right].$ Quasilinear terms
EDQNM type terms

DSS as a subgrid model: An example (Kuan Li)

Hasegawa-Wakatani

$$\frac{\partial \zeta}{\partial t} + \{\phi, \zeta\} = \alpha(\phi - n) - \mu \nabla^4 \zeta \qquad \qquad \kappa = -\partial/\partial x \ln n_0$$
$$\frac{\partial n}{\partial t} + \{\phi, n\} = \alpha(\phi - n) - \kappa \frac{\partial \phi}{\partial y} - \mu \nabla^4 n \qquad \qquad \nabla_{\perp}^2 \phi = \zeta.$$

zonal :
$$\langle f \rangle \equiv \frac{1}{L_y} \int f dy$$
, nonzonal : $\tilde{f} \equiv f - \langle f \rangle$

(Modified) Hasegawa-Wakatani $\frac{\partial \zeta}{\partial t} + \{\phi, \zeta\} = \alpha(\tilde{\phi} - \tilde{n}) - \mu \nabla^4 \zeta$ $\frac{\partial n}{\partial t} + \{\phi, n\} = \alpha(\tilde{\phi} - \tilde{n}) - \kappa \frac{\partial \phi}{\partial y} - \mu \nabla^4 n$

M. Wakatani and A. Hasegawa, "A collisional drift wave description of plasma edge turbulence", Physics of Fluids, 27 (3), 1984.

A Hasegawa and M Wakatani, "Self-Organization of Electrostatic Turbulence in a Cylindrical Plasma", Physical Review Letters, 59 (14), 1987.

Numata, R., Ball, R., & Dewar, R. L, "Bifurcation in electrostatic resistive drift wave turbulence". Physics of Plasmas, 14 (10), 102312, 2007

DSS as a subgrid model: An example (Kuan Li)



 $\kappa = 1, \mu = 10^{-4}, \alpha = 1$

Code verified against simulations with Gkeyll http://ammar-hakim.org/sj/je/je17/je17-hasegawa-wakatani.html

DSS as a subgrid model: An example (Kuan Li)



Conclusions

- Methods from non-equilibrium statistical mechanics may be useful in determining statistics of geophysical/astrophysical systems that include interactions of mean flows and fields with turbulence. Direct Statistical Simulation.
- One can move beyond quasilinearity by
 - Going to higher order in cumulant expansion
 - Generalising the quasilinear approximation (GQL)
- Significant speed-up can be achieved by using a reduced basis.
- This technique may form the basis for conservative statistical sub-grid models.
- For astrophysical fluids/dynamos this may offer a fast, efficient, conservative, self-consistent mean-field theory